



# Hydrological and economic feasibility of mitigating a stressed coastal aquifer using managed aquifer recharge: a case study of Jamma aquifer, Oman

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**Abstract:** This study explored the hydrological and economic feasibility of managed aquifer recharge (MAR) using tertiary treated wastewater (TWW) to mitigate salinity in the coastal aquifer of Jamma, Oman. A steady-state groundwater flow and transport model, using MODFLOW software, was developed and calibrated. Different managerial scenarios were simulated and the results reveal that the Jamma aquifer will be further deteriorated in the next 20 a if it remains unmanaged. The groundwater table will decline further by more than 3 m on average; and the iso-concentration salinity line of 1500 mg/L will advance 2.7 km inland, which will severely affect the farming activities in the area. However, MAR using TWW when integrated with the management of groundwater abstraction (e.g., using modern irrigation systems to reduce the abstraction rate) becomes hydrologically feasible to augment the aquifer storage and control seawater intrusion, and hence improves the farming activities. The results indicate that: (1) injecting TWW in the vicinity of irrigation wells (Scenario A2); (2) investing in smart water meters and online control of pumping from the wells to reduce the abstraction rate by 25% (Scenario B); and (3) a combination of both (Scenario B2) are feasible scenarios with positive net present values. Recharge in upstream areas is found not economically feasible because of the very high investment cost of the installation of pipes to transport the TWW over a distance of 12.5 km. Because of securing funds are challenging, Scenario B would be the best option and the second-best option is Scenario A2. Scenario B2 has the lowest net benefit investment ratio and is very attractive because it entails integrated demand and supply management of groundwater. It is required to reduce pumping and to invest in injecting TWW to improve groundwater quality in the vicinity of irrigation wells and to form a hydrological barrier to control seawater intrusion in the long run.

**Keywords:** managed aquifer recharge; treated wastewater; salinity line; coastal aquifer; Oman

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## 1 Introduction

In arid regions, such as Oman, limited water resources may threaten the development of agriculture,

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industry and tourism. Groundwater is overexploited in densely populated/developed coastal areas, which depletes aquifer storage and causes seawater intrusion that degrades agricultural land and hence adversely affects the socioeconomic aspects of the farming community (MRMWR, 2005). Along with the proper management of water resources, additional water sources are needed to augment stressed coastal aquifers, among which are construction of recharge dams, control/rationing of the groundwater abstraction, and desalination of seawater. The rainfall pattern is irregular, unpredictable and of low rate in Oman (annual average rainfall is approximately 100 mm). Thus, the role of recharge dams is tightly linked to weather conditions, which are characterized by long dry spells and low frequent rainfall (Sen, 2008). Desalinated water is expensive (approximately 1 USD/m<sup>3</sup> as estimated by Zekri et al. (2013)) and has negative environmental impacts. Recently, managed aquifer recharge (MAR) using tertiary treated wastewater (TWW) as a non-conventional water resource has been viewed as a feasible option (Asano and Cotruvo, 2004; Khan et al., 2008; Al-Assa'd and Abdulla, 2010; Missimer et al., 2012; Ebrahim et al., 2015). MAR is the intentional recharge of water into an aquifer either by injection or infiltration and recovery by planned extraction (Hayder Consulting, 2006). MAR has been identified as a potential major water management practice to support groundwater storage in arid and semi-arid areas (Ebrahim et al., 2015).

Abiye et al. (2009) found that TWW can be used for MAR rather than discharge into the Little Akaki River in Ethiopia for storage in the groundwater reservoir because the treated effluent is of better quality than the river water. Li et al. (2006) numerically investigated the impact of MAR on Perth Basin, Western Australia and developed a model which can help to evaluate the effects of different ways of MAR in augmenting groundwater storage and improving the aquifer water quality, particularly along the coast and river margins. MAR also improved the groundwater quality and augmented the storage of an alluvial wadi aquifer in Saudi Arabia (Missimer et al., 2012). Moreover, the injected TWW acted as a hydraulic barrier to decelerate seawater intrusion in the Damour and Jieh regions in Lebanon (Masciopinto, 2013).

Economically, MAR often provides the cheapest water supply (Dillon, 2005; Zekri et al., 2013). Khan et al. (2008) suggested that underground water storage facilities are cheaper than those of surface water storage, with less environmental consequences and evaporation losses. Zekri et al. (2013) proposed the reuse of TWW to farmers through injection/recovery systems via aquifers. The injection of TWW into aquifers for later use in irrigation will further result in natural treatment and quality improvement. Zekri et al. (2013) estimated the current TWW cost in Oman was 0.550 USD/m<sup>3</sup>, while the cost of injection and recovery of TWW in an aquifer is approximately 0.026 USD/m<sup>3</sup>. The quality of water used for aquifer recharge could alter the physical and chemical characteristics of the porous medium. The injected TWW must undergo treatment processes to satisfy drinking water standards and reduce the associated risks (Hayder Consulting, 2006; Dillon et al., 2009). In Oman, Haya Water Company is the main entity responsible for the collection and treatment of sewage water to produce tertiary TWW. The hydrochemistry of MAR practices is beyond the scope of this paper and will be considered in future studies. The volume of tertiary TWW is  $1.00 \times 10^5$  m<sup>3</sup>/d in 2015 (Zekri et al., 2013) and is expected to reach  $2.74 \times 10^5$  m<sup>3</sup>/d by 2030 in Muscat, the capital city of Oman (Zekri et al., 2016).

The objective of this study is to numerically assess the feasibility of MAR using tertiary TWW to mitigate a deteriorated coastal unconfined alluvium aquifer at the Jamma site in the Al-Batinah area of northern Oman.

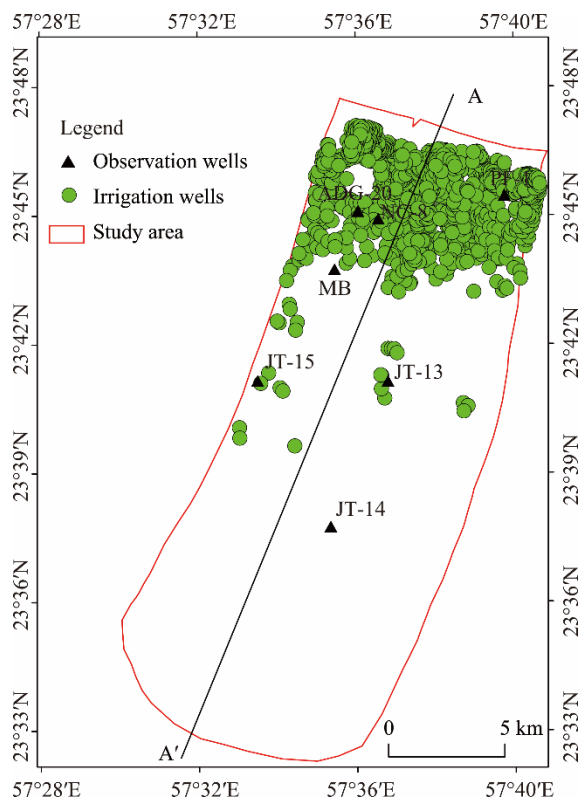
Currently, the Jamma aquifer is mainly used for irrigation purposes. The Al-Batinah plain has approximately 50% of the total agricultural land in Oman (Oman Salinity Strategy, 2012). Over-pumping during the last 45 a has resulted in a significant decline in the groundwater table (up to 5 m), which consequently caused seawater intrusion and soil salinization (MRMWR, 2005; Al Barwani and Helmi, 2006; Zekri, 2008). As a result, suitable agricultural lands in the Al-Batinah plain reduced by 7% in 2000–2005, with negative socioeconomic consequences (Al Barwani and Helmi, 2006; Zekri, 2009; Oman Salinity Strategy, 2012). The injection of TWW into the aquifer will form a hydraulic barrier against seawater intrusion, improve water quality within the vicinity of the injection and furnish additional water for irrigation. In this study, the clogging effect during

the artificial recharge as discussed by Voudouris (2011) is not considered at this stage as the treatment level is tertiary with additional purification steps (in fact Quaternary).

## 2 Materials and method

### 2.1 Study area

The Jamma aquifer (23°35'53"N, 57°35'14"E) is an important alluvium unconfined aquifer (Al Barwani and Helmi, 2006) that is located in the southeast part of the Al-Batinah coast in Oman (Wadi Al-Fara catchment) and covers an area of approximate 295 km<sup>2</sup> (Fig. 1). The study area, Jamma, is characterized by warm winter with low humidity, very hot and humid summer, along with low and irregular rainfall. The annual mean rainfall is approximately 60 mm, with an annual potential evapotranspiration (ET) of 1644 mm (MWR, 1996). The highest annual rainfall (240 mm) is recorded in the southern region of the study area (above 3000 m a.s.l.) compared with the coastal region (80 mm) from which the major recharge to the coastal aquifer occurs (Weyhenmeyer et al., 2000). The mean annual temperature is 28.5°C. There are approximately 1037 unmonitored/unregulated irrigation wells clustered along the coastal strip with a width of 5 km (Fig. 1) that mainly tap the upper part of the Quaternary alluvium layer because the vast majority of the wells are shallow-dug wells. The groundwater abstraction records in the study area are very limited. The total groundwater abstraction from irrigation wells has been estimated based on field surveys and measurements conducted by the Ministry of Regional Municipalities and Water Resources (MRMWR) of Oman. The total volume of irrigation water in the study area was estimated based on the irrigation water requirement per hectare under different irrigation systems (Table 1; Zekri et al., 2016). The cropped area is approximately 10.9 km<sup>2</sup>, with 80% of it under traditional flood irrigation and the rest under modern irrigation systems (MRMWR, 2005). The total abstraction from irrigation wells is approximately  $89 \times 10^6$  m<sup>3</sup>/a, with 235 m<sup>3</sup>/d as an average abstraction rate per well. Irrigation returns to the aquifer (as a source of recharge) is ignored



**Fig. 1** Location of the irrigation and observation wells in the study area

provided the depth to the water table varies from 61.4 m at the southern boundary to 16.0 m in the vicinity of the coastal boundary (average water table, 31.6 m). The groundwater table was monitored using seven observation wells by the MRMWR, which were used in the calibration of the numerical model (Fig. 1).

**Table 1** Irrigation water requirement

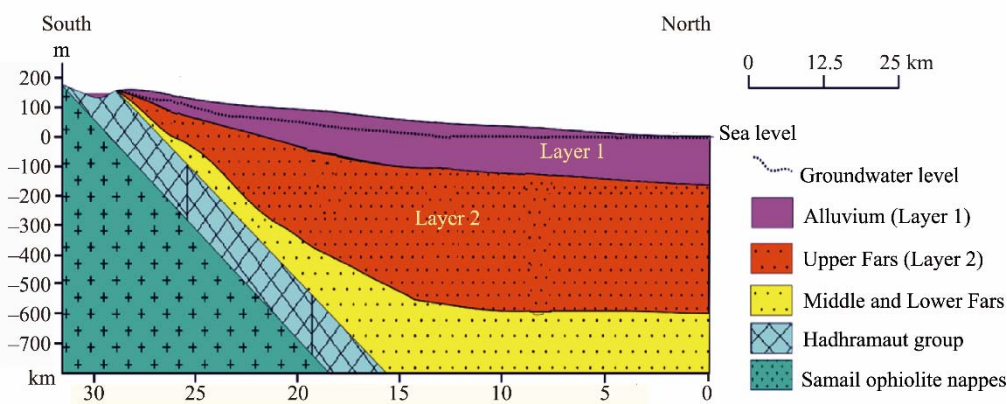
Irrigation system	Irrigation water requirement ( $\times 10^3 \text{ m}^3/(\text{hm}^2 \cdot \text{a})$ )
Flood irrigation	96.469
Modern irrigation	22.917
Modern+flood irrigation	42.012

Note: Modern irrigation includes drip and sprinkler irrigations.

## 2.2 Geology and hydrogeology of the study area

The geology of Wadi Al-Fara catchment was previously investigated by many researchers (e.g., Searle and Malpas, 1980; Stanger, 1986; Lakey et al., 1995). The Wadi Al-Fara catchment comprises two distinct geomorphological zones. The first zone (southern part) is mountainous/piedmont upper catchment dominated by Samail Nappes, as well as Hadhramaut Group (HG) sedimentary rocks. The second (northern part, Jamma area) is alluvial fans, which is a plain that extends through the central and lower reaches of the catchment to the coast. The aquifer is unconfined and modeled as a two-layered system, as shown in the geological cross-section (Fig. 2). Layer 1 (quaternary alluvium) has a low clay content (with a hydraulic conductivity of 8–70 m/d) whereas layer 2 (tertiary alluvium, or as named locally "Upper Fars Formation") is more compacted and characterized with lower hydraulic conductivity (0.1–20.0 m/d). Those values of hydraulic conductivities were used in the calibration process (Table 2). The thickness of layer 1 ranges from 80 to 220 m. The thickness of layer 2 varies between 195 and 430 m (Fig. 2). The specific yield varies between 0.01 to 0.20 and the porosity is equal to 0.25.

The bottom elevation of the alluvial unconfined aquifer in the study area was obtained from geophysical data and lithological descriptions of piezometers (MAF, 1986; MRMWR, 2006), while the bottom elevation of the Upper Fars aquifer in the study area was extrapolated from geological cross-sections of piezometers located close to the boundary with a neighboring catchment (Wadi Bani Ghfir) (MRMWR, 2006).



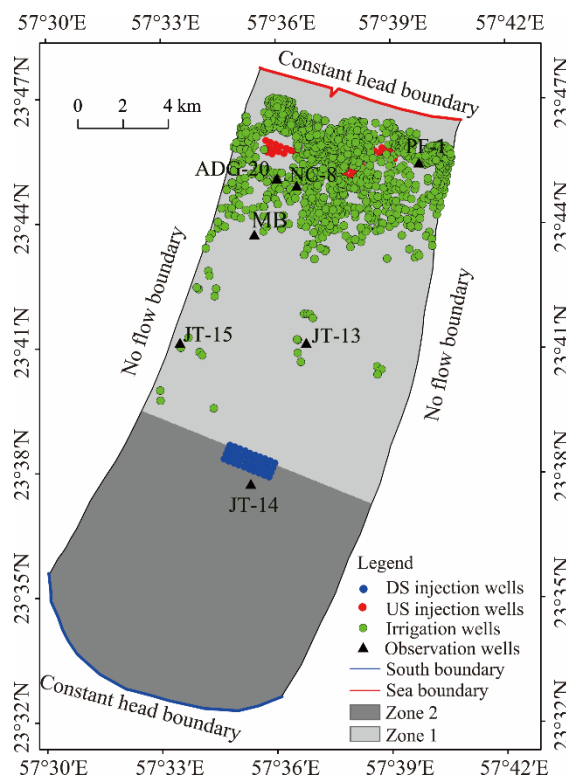
**Fig. 2** Geological cross-section of the study area, Jamma aquifer, in the Wadi Al-Fara catchment

## 2.3 Numerical model setup

A three-dimensional finite-difference groundwater flow model MODFLOW-2005 (Harbaugh, 2005) was used to simulate groundwater flow using ModelMuse (Winston, 2009) as the graphical user interface (GUI). The model was discretized by a 30 m $\times$ 30 m grid, resulting in a domain of 952 rows and 610 columns and 2 layers with 1,161,440 cells (2 $\times$ 952 $\times$ 610), of which 655,390 are active cells.

The southern boundaries hydraulic conductivity zones and injection wells are shown in Figure 3. Constant head boundaries were assigned to the coastline (water head, 0 m) and southern boundary (water head, 95.0 m as estimated from piezometric map) to account for the deep percolation from the mountainous area that receives a higher amount of rainfall (240 mm/a) and is considered as the header tank for the coastal area (Fig. 2). Analysis of head hydrograph data for the observation well, located at the upstream boundary, supports the assumption of a "constant head boundary" as the water table remains constant over the last 15 a. No flow boundaries were assigned to the eastern and western boundaries of the model domain since the regional flow is normal to the coastline. The direct recharge from rainfall in the coastal plain area was assumed to be 20% of the annual rainfall (Lakey et al., 1995; Walther et al., 2014; Ebrahim et al., 2015). The potential ET of 1644 mm/a was assigned with an extinction depth of 12 m (Christmann and Sonntag, 1987; and Abdalla, 2008). The processing of the pumping test data divided the modeled area into two main zones (Fig. 3): the upstream part Zone 2 had a lower hydraulic conductivity of 0.1–8.0 m/d and the downstream Zone 1 had a higher hydraulic conductivity of 10.0–50.0 m/d.

The seawater salinity concentrations in the Oman Sea range from 40,000 to 48,000 mg/L (Greenlee et al., 2009). In this study, a constant concentration of 40,000 mg/L was assigned along the coastal boundary, with an assumed dispersion coefficient of 200 m<sup>2</sup>/d. The solute transport for seawater intrusion was simulated based on constant density approach using the MT3DMS code (Zheng and Wang, 1999). The density effect was ignored in this study as the interest is a regional scale and the effects of the density on the flow and transport processes of saline water could be neglected for the purpose of this study. The simulation period is 20 a, with 244 stress periods of 30 d each. The 20-a period is selected to explore the aquifer responses over a relatively short hydrological time to MAR practice. TWW was injected during four months winter (during which an excess of TWW is available, as per the TWW industry), followed by eight months of recovery or without injection (WMAR).



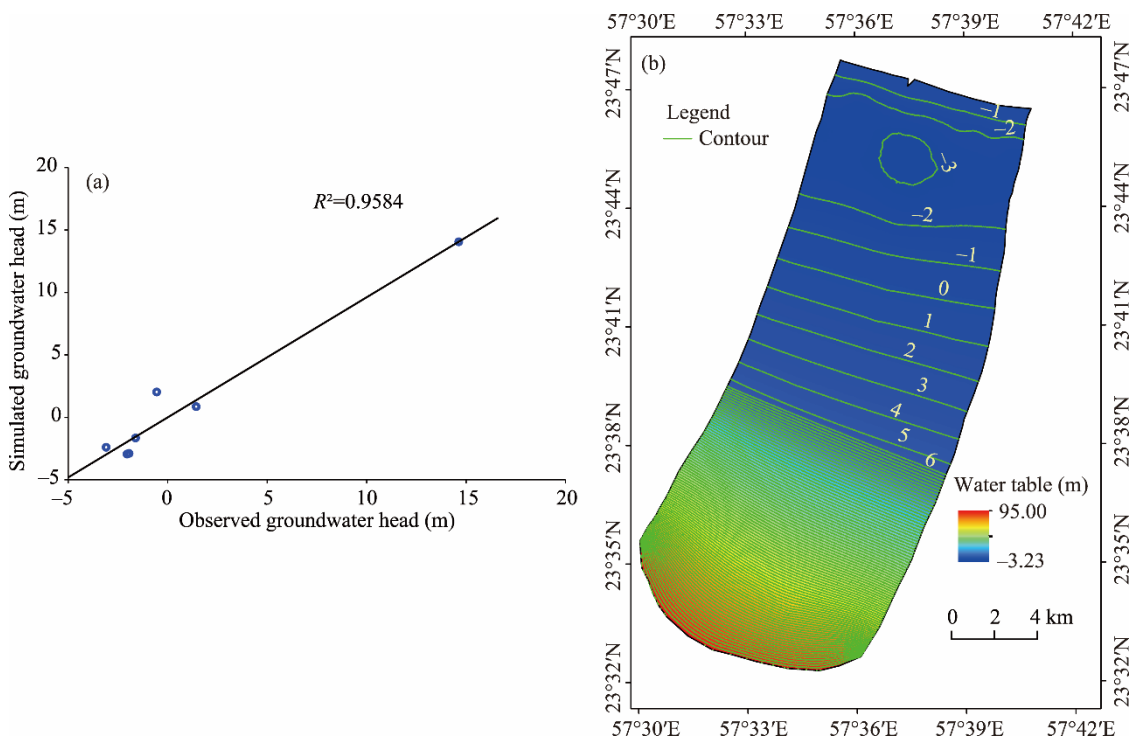
**Fig. 3** Locations of irrigation (green dot) and observation (black triangle) wells, the sea and southern boundaries, hydraulic conductivity zones and locations of forty upstream (US) and downstream (DS) treated wastewater injection wells

## 2.4 Sensitivity analysis and calibration

A sensitivity analysis was carried out using UCODE 2005 (Poeter et al., 2005) with the help of ModelMate (Banta, 2011). The lack of time series data for irrigation wells (private wells) is a challenge that hindered the transient state calibration process. The model is calibrated for the steady state using data from 2005 from the seven available piezometers (Fig. 3). The values of the calibrated parameters are presented in Table 2. The results of the calibrated groundwater model shows a correlation factor of 0.96 (Fig. 4a). The simulated head distribution using the calibrated parameters is presented in Figure 4b. The mean error is  $-0.04$  m, mean absolute error is  $0.90$  m, and root mean square error is  $1.17$  m, which shows that the model performs well.

**Table 2** Calibrated model parameters for the Jamma aquifer

Parameter	Value (m/d)	Description
HKL1Z1	35.0	Hydraulic conductivity of Zone 1 in layer 1
HKL1Z2	8.0	Hydraulic conductivity of Zone 2 in layer 1
HKL2Z1	20.0	Hydraulic conductivity of Zone 1 in layer 2
HKL2Z2	0.3	Hydraulic conductivity of Zone 2 in layer 2
Rech	$1.80247\text{E-}06$	Recharge from precipitation
Extinction depth	12 m	Evapotranspiration extinction depth



**Fig. 4** Simulated versus observed groundwater heads for the steady state using seven monitoring wells at the Jamma site (a) and simulated groundwater head using the calibrated parameters (b)

The water budget of the calibrated model is summarized in Table 3. Using ZONEBUDGET program (Harbaugh, 1990) as a post-processor step, the inflow from the sea boundary into the aquifer was calculated. The total amount of inflow from the coastal boundary into the aquifer is  $171.702 \times 10^3$  m<sup>3</sup>/d, while the recharge through the southern boundary is  $79.288 \times 10^3$  m<sup>3</sup>/d and the direct recharge from rainfall to the aquifer is approximately  $0.534 \times 10^3$  m<sup>3</sup>/d, which is 0.2% of the daily inflow amount. The abstracted volume by irrigation wells is  $2.43.695 \times 10^3$  m<sup>3</sup>/d, which is 97% of the total outflow rate, while the outflow by ET is estimated at  $7.829 \times 10^3$  m<sup>3</sup>/d, which accounts for 3% of the total outflow budget from the modeled area and is an acceptable volume given the

depth of the water table (16.0–61.4 m from downstream to upstream). The over abstraction of groundwater resulted in a water depression in the vicinity of the irrigation wells, reaching approximately 3.3 m, causing seawater intrusion in the area.

**Table 3** Water balance of the steady state calibrated model

	Inflow		Outflow		Difference
	( $\times 10^3$ m <sup>3</sup> /d)	(%)	( $\times 10^3$ m <sup>3</sup> /d)	(%)	
From the coastal boundary (Oman Sea)	171.702	68.3	0	0	171.702
inflow from the south boundary	79.288	31.5	0	0	79.288
Pumping wells	0	0	243.695	97	-243.695
ET	0	0	7.829	3	-7.829
Recharge	0.534	0.2	0	0	0.534
Total	251.524	100	251.524	100	0

## 2.5 Simulated scenarios

Eight scenarios (which are divided into 3 main groups: A, B and C) were simulated and are presented in Table 4. The simulation time is set as 20 a, with 244 stress periods of 30 d each. Scenario A considers the case "business as usual", which simulates the current situation assuming that no changes (in terms of management and climatic conditions) will take place for the next 20 a. Scenario A will thus be considered as the base scenario, with which the results of the other scenarios are compared. The responses of the aquifer to MAR using TWW in two injection locations were simulated (scenarios A1 and A2). The locations of the injection wells were taken into consideration of land availability because the zone is densely populated and privately owned. The locations of wells are given as Zone 1 for upstream and Zone 2 for downstream (vicinity of farms) in Figure 3. The injected volume is based on the availability of excess TWW ( $31 \times 10^6$  m<sup>3</sup>/a), as reported by the Haya Water Company (Zekri et al., 2013). Scenarios B and C are based on policies recommended by the Ministry of Agriculture and Fisheries of Oman (Oman Salinity Strategy, 2012) and by the MRMWR in Oman (Abdel-Rahman and Abdel-Magid, 1993), as shown in Table 4.

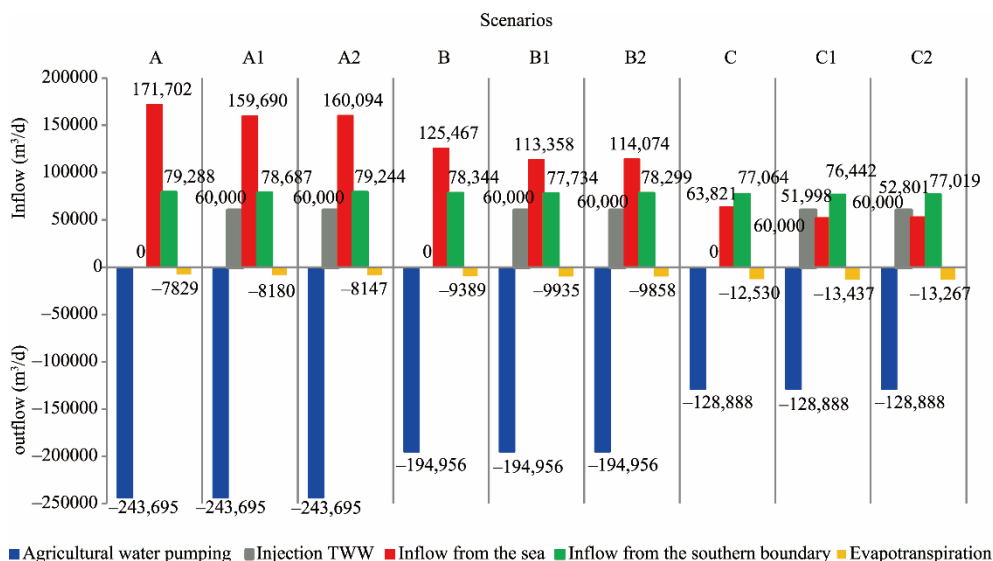
**Table 4** Description of the simulated scenarios

Scenario	Location of the injection	Description	Injection rate ( $\times 10^3$ m <sup>3</sup> /d)	Number of wells	Period of injection	Abstraction rate ( $\times 10^3$ m <sup>3</sup> /d)
A	NA	Represents the current situation	NA	NA	NA	243.695
A1	Upstream	To test the feasibility of upstream injection on augmentation of aquifer storage	1.500	40	4-mth/8-mth	243.695
A2	Downstream	To test the feasibility of MAR to form a hydraulic barrier against seawater intrusion	1.500	40	4-mth/8-mth	243.695
B	NA	Abstraction volume reduced by 25%	NA	NA	NA	194.956
B1	Upstream	Abstraction volume reduced by 25%	1.500	40	4-mth/8-mth	194.956
B2	Downstream	Abstraction volume reduced by 25%	1.500	40	4-mth/8-mth	194.956
C	NA	Abstraction volume reduced by 50%*	NA	NA	None	128.888
C1	Upstream	Abstraction volume reduced by 50%	1.500	40	4-mth/8-mth	128.888
C2	Downstream	Abstraction volume reduced by 50%	1.500	40	4-mth/8-mth	128.888

Note: A, the base scenario; NA, not applicable; MAR, managed aquifer recharge; Upstream, the injection was performed 16.5 km upstream of the coastline (approximately 12.5 km from the water supply line); Downstream, in the vicinity of the farming area and approximately 2.7 km from the water supply line; 4-mth/8-mth, 4 months MAR followed by 8 months WMAR (without managed aquifer recharge); \*, data source: Abdel-Rahman and Abdel-Magid, 1993.

### 3 Results and discussion

The effects of the different simulated scenarios on mitigating the aquifer were analyzed at the end of the simulated period. With respect to the base scenario A, the injection of  $60.000 \times 10^3 \text{ m}^3/\text{d}$  of TWW in the upstream location (Scenario A1) causes a decrease in the inflow of water through the coastline ( $12.012 \times 10^3 \text{ m}^3/\text{d}$ ) and the southern boundaries ( $601 \text{ m}^3/\text{d}$ ) (Fig. 5). The change in water volume discharge via ET is small. The groundwater table rises by 0.85 m on average.



**Fig. 5** Water balance of the different scenarios at WMAR20 (end of 240 stress periods) for the simulated scenarios

Moving the injection wells to the vicinity of the farms (the stressed part of the aquifer, Scenario A2) due to the injection of  $60.000 \times 10^3 \text{ m}^3/\text{d}$  of TWW in the downstream part of the aquifer, the inflows from both the coastline and southern boundaries into the aquifer decreased by  $11.608 \times 10^3$  and  $44 \text{ m}^3/\text{d}$ , respectively, against to Scenario A. It is similar to Scenario A1 because of the effects of the active irrigation wells in simultaneously recovering the injected water and hence the slow gain in storage. The groundwater table rises by 0.2 m on average (Table 5).

In Scenario B, the reduction in the volumetric abstraction rate was 25% ( $194.956 \times 10^3 \text{ m}^3/\text{d}$ ) and there was no MAR practice. The inflow from both the coastline and the southern boundaries decreased by  $45.200 \times 10^3$  and  $944 \text{ m}^3/\text{d}$ , respectively, with respect to Scenario A. The impact of reducing the abstraction is much higher than that induced by MAR in the scenarios A1 and A2. The recovery in the water table is 0.82 m but a great volume of water is lost through ET (17%).

When the reduction in abstraction from irrigation wells in Scenario B is complemented with MAR at the locations of the scenarios B1 and B2, the inflows from both the coastline and southern boundaries decreased by  $48.300 \times 10^3$  and  $1554 \text{ m}^3/\text{d}$ , respectively, for scenario B1 compared to Scenario A. The groundwater table rose 1.66 m on average (Table 5). For Scenario B2, the aquifer receives approximately  $57.628 \times 10^3 \text{ m}^3/\text{d}$  of seawater, which represents a reduction of 33.5% compared to Scenario A. Injection at the downstream site in Scenario B2 reduces seawater intrusion by only 5% compared to Scenario B1 and by 7% with respect to Scenario B. MAR increased the average water table by 100% in Scenario B1 and 23% in Scenario B2 with respect to Scenario B. The active role of irrigation wells on the recovery of the injected TWW provided less opportunity for the water table mound to develop and hence the aquifer storage to recover (Table 5).

A further drop in the volumetric abstraction rate of 50% (Scenario C,  $129.000 \times 10^3 \text{ m}^3/\text{d}$ ), even in the absence of MAR, will decrease the seawater intrusion rate by 63% compared to Scenario A (inflow from the sea decreased by  $108.000 \times 10^3 \text{ m}^3/\text{d}$ ). This reduction allows aquifer storage recovery, which is reflected in the increase of the average groundwater table by 1.91 m (Table 5). Obviously, losses via ET increase by 60%, which is relatively small compared to other outflows

from the modeled system. Similar to the scenarios B1 and B2, the injection of TWW will further enhance the hydrological condition of the aquifer. The inflow of saline water decreases by 70% and the average groundwater table rises by 2.75 m (Table 5) when Scenario C1 is simulated. When MAR is shifted to the downstream locations (Scenario C2), the induced changes are almost the same as those of Scenario C1.

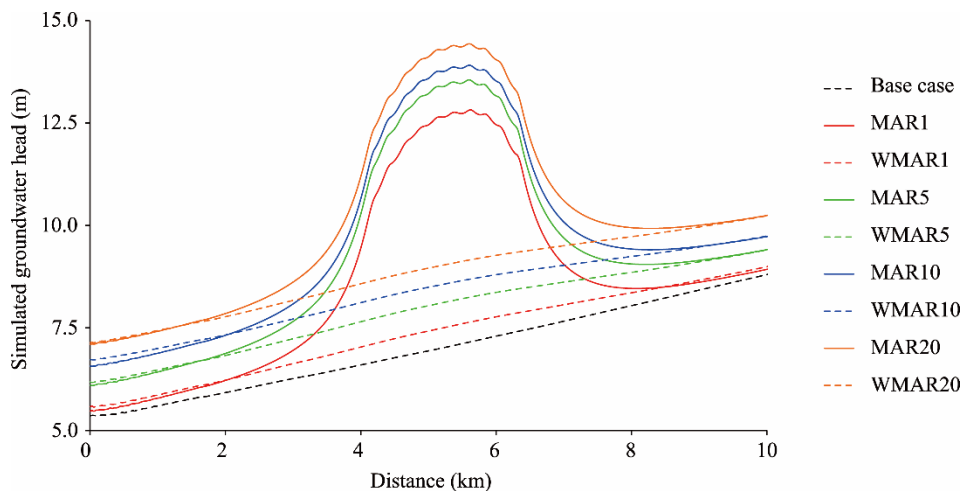
**Table 5** Abstraction and injection rates as well as changes in inflow from the sea boundary, discharge through evapotranspiration, and the changes in the average groundwater table of all scenarios against Scenario A

Scenario	Abstraction from agricultural wells	Injection rate	Change in inflow from the sea boundary ( $\times 10^3 \text{ m}^3/\text{d}$ )	Change in evapotranspiration	Average change in groundwater table (m)
A	243.695	0	NA	NA	NA
A1	243.695	60.000	12.012	0.351	0.85
A2	243.695	60.000	11.608	0.318	0.20
B	194.956	0	46.235	1.560	0.82
B1	194.956	60.000	58.344	2.106	1.66
B2	194.956	60.000	57.628	2.029	1.01
C	128.888	0	107.881	4.701	1.91
C1	128.888	60.000	119.704	5.608	2.75
C2	128.888	60.000	118.901	5.438	2.10

Note: NA, not applicable.

The results suggest that MAR is more efficient in restoring the stressed coastal aquifer when integrated with the management of abstraction from irrigation wells. It is obvious that injecting more TWW will augment the aquifer storage, but the availability of TWW is a limitation. Moreover, the cost-benefit aspect must be considered when the optimization of MAR and irrigation practices are considered.

The developed groundwater mound for a selected scenario is plotted against time at a cross-section (which passes through the zone of the injection wells) across the modeled area (extending between the eastern and western boundaries; Fig. 3) for the selected stress periods. A detailed analysis of the results shows that Scenario A1 creates a mound of 8.28 m after MAR20 (at end of MAR duration for 20 a). The mound dissipates gradually when injection ceases and extends over a long distance of approximately 3.5 km from the center of the zone of the injection wells. MAR enhances the aquifer storage, as reflected in the gradual rise in the mound after successive and consistent injections (Fig. 6). The buildup of a hydraulic head over time due to MAR will help to

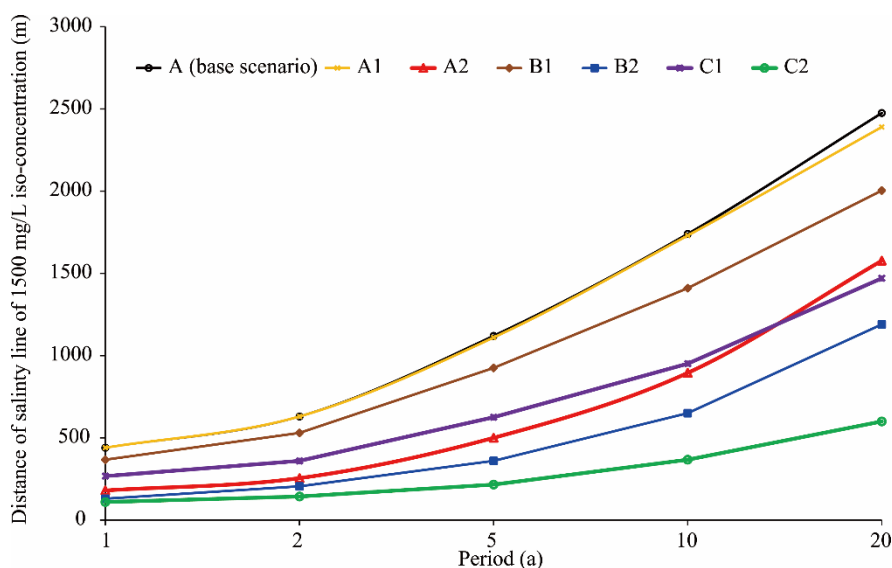


**Fig. 6** Simulated groundwater heads at a line crossing the injection wells at the end of different durations of MAR and WMAR (MAR1, WMAR1, MAR2, WMAR2, MAR5, WMAR5, MAR10, WMAR10, MAR20 and WMAR20) for Scenario A1 with respect to the base scenario A

reduce the inflow from the sea and hence improve and sustain farming activities. However, the base scenario suggests that the water table will further drop by more than 3 m in average in case the aquifer remains unmanaged.

The 1500 mg/L iso-concentration line for Scenario A and scenarios A1, A2, B1, B2, C1 and C2 are considered (Fig. 7) to observe the effects of MAR on the recession of the saline and fresh water interface. The salinity line recedes in the seaward direction by nearly 1.0 km at the end of 20 a for Scenario A2 and by approximately 1.3 km for Scenario B2. Figure 6 clearly shows that MAR using TWW acts as a hydraulic barrier against seawater intrusion. It is, however, less effective in this regard when the injection is implemented far from the coastline in the upstream area. However, the upstream injection is expected to further purify the injected water when it is filtered out by the subsurface porous domain, and perhaps its influence on the hydrologic system will be observed in a period longer than the simulated periods in this research.

The results show that MAR becomes more effective in both restoring the head distribution in the depressed zone and augmenting the groundwater quality when MAR is applied in the downstream areas and integrated with a 50% reduction in abstracted irrigation water (Scenario C2).



**Fig. 7** Salinity line of 1500 mg/L iso-concentration for the scenarios A, A1, A2, B1, B2, C1, and C2

Table 6 presents the salinized water volume and its changes considering the 1500 mg/L iso-concentration salinity line for Scenario A and simulated scenarios at the end of the simulation time. Considering the 244<sup>th</sup> stress period, a volume of  $2.561 \times 10^9$  m<sup>3</sup> is salinized for Scenario A. In Scenario C2, 79% of the salinized water is cleaned to a salinity level lower than 1500 mg/L and is hence becomes suitable for farming activities (for the types of crops and trees cultivated in the study area, i.e., date palm trees, lime trees and fodder crops). In the absence of aquifer management (e.g. MAR and efficient irrigation systems), the saline water interface is expected to intrude a distance of 2.7 km landward direction.

A cost-benefit analysis was applied to evaluate the economic value of the scenarios using a 20-a life span and 5% discount rate. All the scenarios that involve recharge in the upstream area are not economically feasible due to the very high investment cost to transport water by 16.5 km. Scenario B does not involve any recharge to the aquifer and the investment cost is related to the cost of installation of smart groundwater meters in each of the 1037 wells and online monitoring of the pumping (Zekri et al., 2017). Scenario B2 is a combination of scenarios A2 and B. It assumes that the injection of treated wastewater will take place in the downstream area, and at the same time, pumping will be controlled as in Scenario B.

Given that financial resources for investments are scarce, the decision criteria on the best scenario should rely on the net benefit investment ratio. The highest ratio of 4.41 corresponds to

Scenario B. Even though Scenario B2 has the lowest net benefit investment ratio, it is very attractive from a social perspective because it involves two measures at a time. Thus, farmers are requested to cut pumping and the government will invest in recharge to improve the quality of the groundwater. The joint effort of the ministry and farmers will provide a higher chance of success than acting from a single side only.

**Table 6** Salinized water volume for different simulated scenarios considering the 1500 mg/L iso-concentration salinity line at the end of the simulation period

Scenario	Salinized water volume ( $\times 10^9 \text{ m}^3$ )
A	2.561
A1	2.488
A2	1.443
B	2.123
B1	2.030
B2	1.067
C	1.400
C1	1.300
C2	0.537

## 4 Conclusions

Jamma aquifer will further deteriorate if the current practices continue without a plan for groundwater management for the next 20 a, which will lead to more than 3.0 m average decline in the groundwater table and 2.7 km further intrusion in the 1500 mg/L salinity line. By contrast, using MAR will result in raising the groundwater table by 0.85 m and cause the salinity line to recede by more than 1000 m. The best managerial results can be achieved when the management of groundwater abstraction (e.g., using modern irrigation systems) is integrated with MAR using TWW. However, the scenarios with very high investment and maintenance costs are not economically feasible. The main challenge in practicing MAR for all aquifers along Al-Batinah coast is the availability of TWW. The numerical simulations show that MAR can be applied for a single aquifer for a certain period of time until the aquifer is cured and can then be moved to another aquifer. As a feasible management practice to augment water resources in salinized coastal aquifers in arid areas, MAR can help improve farming profitability and sustainability as well as food security in Oman.

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